IDEA AND PERSPECTIVE

Predictability in community dynamics

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Abstract
The coupling between community composition and climate change spans a gradient from no lags to strong lags. The no-lag hypothesis is the foundation of many ecophysiological models, correlative species distribution modelling and climate reconstruction approaches. Simple lag hypotheses have become prominent in disequilibrium ecology, proposing that communities track climate change following a fixed function or with a time delay. However, more complex dynamics are possible and may lead to memory effects and alternate unstable states. We develop graphical and analytic methods for assessing these scenarios and show that these dynamics can appear in even simple models. The overall implications are that (1) complex community dynamics may be common and (2) detailed knowledge of past climate change and community states will often be necessary yet sometimes insufficient to make predictions of a community's future state.

Keywords
Alternate states, chaos, climate change, community assembly, community climate, community response diagram, disequilibrium, hysteresis, lag, memory effects.

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INTRODUCTION

Understanding how communities respond to climate change is necessary for predictive modelling of global change and for identifying the processes that have shaped contemporary biodiversity patterns. A key aspect is the degree of lag in the response of community composition to contemporary climate conditions. By lag, we mean the *amount* the community is out of equilibrium with the observed climate, in either a positive or negative direction. The equilibrium no-lag state of a community should reflect a set of species with climate niche optima close to the observed climate at a given location. However, since climates change over time, a range of transient disequilibrium community states could be achieved, in which the community's composition is lagged relative to climate.

There are two extreme hypotheses for the magnitude of lags in the response of community composition to climate change. No-lag responses are thought to occur when species respond through local persistence via high niche plasticity or niche adaptation, or rapid extinction at trailing range edges (Hampe & Petit 2005) and/or efficient long-distance dispersal and range expansions at leading range edges. In this case, the community responds instantly to climate change and is in an equilibrium state. Conversely, lagged responses are thought to occur when species have limited dispersal ability, have long persistence times or when the regional pool does not include more appropriate species (Svenning & Sandel 2013; Blonder *et al*. 2015). In this case, the community is in a transient disequilibrium state that will change both when the climate varies and when the climate does not vary. These two ideas form the conceptual foundation for several large bodies of work and are thought to encompass the range of possible community responses to climate change (Ackerly 2003), with the speed and type of species response of fundamental interest for predictive modelling and for biodiversity conservation (Nicotra *et al*. 2010; Hoffmann & Sgro 2011; La Sorte & Jetz 2012).

The no-lag hypothesis proposes that at a given time the species composition of a community is in equilibrium with the observed climate at that location, assuming that an equilibrium can be defined over the temporal or spatial scale of interest (Svenning *et al*. 2015). That is, the species found in a community will have climate niches that are close to the observed climate. The implication is that, in a new climate, species with well-matched niches that are already present will persist, other species with well-matched niches will rapidly immigrate and become present, and species with poorly matched niches will rapidly die and become absent. This hypothesis is implicit in many decades of work assuming that vegetation–climate associations represent consistent physiological responses to environment (von Humboldt & Bonpland 1807; Whittaker 1967) and that have often been used to reconstruct climate from paleoecological evidence for pollen,

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chironomids, diatoms, etc., based on transfer functions (Guiot *et al.* 1989; Gasse *et al.* 1995; Brooks & Birks 2000), coexistence intervals (Mosbrugger & Utescher 1997; Pross *et al.* 2000) or probability densities (Kühl *et al.* 2002). Many of these climate reconstruction approaches assume that species-environment relationships are constant and instantaneous, without considering the consequences of this assumption. The no-lag hypothesis is also implicit in the vast majority of environmental niche modelling/species distribution modelling studies that predict climate change responses (Birks *et al.* 2010; Peterson 2011). This hypothesis is a simple baseline assumption that finds support at multiple scales (e.g. both continental extents over sub-millennial to millennial timescales [e.g. in multi-taxon responses to Younger Dryas climate changes in Switzerland (Birks & Ammann 2000) or across the late Quaternary in North America (Shuman *et al.* 2009; Williams *et al.* 2011)] and is consistent with many species having niches that are well predicted by their range limits (Lee-Yaw *et al.* 2016). However, the hypothesis has also been criticised. One major issue is that its assumption of very fast species response can be unrealistic (Campbell & McAndrews 1993; Guisan & Thuiller 2005; Araújo & Peterson 2012). Another important issue is that realised niches may shift relative to the observed climate due to changes in the available climate space or in biotic interactions (La Sorte & Jetz 2012; Veloz *et al.* 2012; Maiorano *et al.* 2013). As such, the realised niche of a species may be a poor proxy for the fundamental niche and may not necessarily be matched to the observed climate (Jackson & Overpeck 2000; Jordan 2011).

Alternatively, lag hypotheses argue that the species composition of a community at a given time is in disequilibrium with contemporary climate (Svenning & Sandel 2013; Blonder *et al.* 2015). That is, the species found in a community may be poorly suited to the climates at the site, despite other species not occurring in the community having better-suited climate niches (Davis 1984; Webb 1986; Dullinger *et al.* 2012). Proposed mechanisms include resident species persisting via survival of long-lived individuals (Eriksson 1996; Holt 2009; Jackson & Sax 2010), species interactions producing microscale conditions that remain favourable (Schöb *et al.* 2012; De Frenne *et al.* 2013), or no immigration of more appropriate species because of priority effects (Fukami *et al.* 2005, 2010), dispersal limitation (Svenning & Skov 2007) or species absence from the regional pool (Blonder *et al.* 2015). These processes together would produce a lag between communities' composition and climate. This hypothesis is reflected in a broad literature showing vegetation lag to climate in forests in the Americas (Webb 1986; Campbell & McAndrews 1993; Blonder *et al.* 2015) and in Europe (Birks & Birks 2008; Bertrand *et al.* 2011; Normand *et al.* 2011; Seddon *et al.* 2015), in bird communities (DeVictor *et al.* 2008), as well as in a range of other paleoecological data (reviewed in Davis (1981) and Svenning *et al.* (2015)).

Here, we argue that there is not a dichotomy between lag and no-lag hypotheses. Rather, there is a continuum of lag hypotheses that encompasses more scenarios than have been previously considered. We show that a broader set of possibilities can lead to unintuitive or difficult-to-predict community responses. We then provide a set of quantitative tools for detecting these scenarios in empirical data. Lastly, we demonstrate that simple models of community processes can generate all of these scenarios.

COMMUNITY RESPONSE DIAGRAMS AS DIAGNOSTICS OF DYNAMICS

Lags and lag hypotheses can be measured by comparing a community's composition to the climate conditions in the community. Making these concepts precise requires defining several concepts (Box 1). These concepts are presented and defined for a single climate axis and variable (e.g. temperature). They can be extended to multiple climate axes using vector approaches (Blonder *et al.* 2015), but are illustrated here in a single dimension for clarity.

First, the location of the community has an observed climate, which is given by a function $F(t)$ (Fig. 1a). This variable changes potentially independently from the state of the community and can be measured without knowledge of the community state, for example with a thermometer for a temperature axis.

Second, the community itself has an inferred climate, which is given by a function $C(t)$ (Fig. 1a). This variable reflects the value of the climate along this axis most consistent with the occurrence of all species at time t. It can be calculated by overlapping the fundamental niches of species in the community. For example, a community with cocoa and banana trees would have a warm inferred climate along a temperature axis, while a community with blueberry and snowberry bushes would have a cold inferred climate. Multiple species assemblages might all yield the same value of $C(t)$.

This concept of $C(t)$ is already widely and implicitly used across fields, although using different terminology. It is widely used in multi-taxon paleoclimate reconstructions (ter Braak & Prentice 1988; Guiot *et al.* 1989; Birks *et al.* 2010; Harbert & Nixon 2015). Additionally, it underlies the definitions in community ecology for a community temperature index (DeVictor *et al.* 2008; Lenoir *et al.* 2013), a floristic temperature (De Frenne *et al.* 2013) and a coexistence interval (Mosbrugger & Utescher 1997; Harbert & Nixon 2015).

Third, the community climate lag can be defined as the difference between the observed and the inferred climate (Fig. 1a). This metric has been previously used in several studies of ecological disequilibrium (Davis 1984; Webb 1986; Bertrand *et al.* 2011; Blonder *et al.* 2015). If these two values are closely matched, then the lag is small; alternatively, if they are not closely matched, then the lag is large.

These statistics can be visualised and combined with a community response diagram. This diagram is a time-implicit parametric plot of the observed climate $F(t)$ on the *x*-axis and the inferred climate response $C(t)$ on the *y*-axis (Fig. 1b). Using dynamical systems terminology (Katok & Hasselblatt 1997; Beisner *et al.* 2003), $F(t)$ would be considered a parameter (exogenous to the system) and $C(t)$ would be considered a state variable (endogenous to the system). The diagram is similar to a phase space diagram of dynamical systems research (e.g. Sugihara *et al.* (2012)) that plots multiple state variables as time-implicit curves, but is different in that $F(t)$ is not a state variable. It also is similar to

Box 1 Definitions of lag statistics for community dynamics

Consider a community containing a set of $\{i\} \in (1, k)$ species at time t. Each resident species i has a fundamental niche function that can be described by a relative fitness over a given niche axis. Suppose that each of these niche functions has a modal value of $N_i(t)$.

The location of the community has an observed climate $F(t)$. The inferred climate of the community also can be defined as the mean of the niche optima of all species (Fig. 1):

$$C(t) = E[N_i(t)]. \qquad (B1-1)$$

More sophisticated definitions (e.g. abundance-weighted means or medians across species) are possible and potentially more useful in low-richness communities.

We can also define a measure of uncertainty in the inferred climate, $\sigma(t)$, as the standard deviation of the modal niche values:

$$\sigma(t) = \sqrt{E\left[(N_i(t) - C(t))^2\right]}. \qquad (B1-2)$$

If a community is comprised of species with similar $N_i(t)$ values, then $\sigma(t)$ is close to zero; alternatively, if species have a wide range of $N_i(t)$ values, then $\sigma(t)$ is large. Large values of $\sigma(t)$ can also represent community lag resulting from differences in species responses to changing climatic conditions, but we primarily consider them as uncertainties in the context of empirical data.

The community climate lag can be defined as the difference between the inferred climate and observed climate. It can be calculated at any given time t:

$$\Lambda(t) = C(t) - F(t). \qquad (B1-3)$$

because of linearity, the standard deviation (uncertainty) of $\Lambda(t)$ is also equal to $\sigma(t)$.

The mean absolute deviation can be defined as:

$$\overline{|\Lambda|} = \frac{1}{t_{\max}} \int_0^{t_{\max}} |\Lambda(t)| dt, \qquad (B1-4)$$

where generally the statistic would be calculated for $t_{\max} \to \infty$.

The maximum state number can be defined as the largest number of real values of C corresponding to any of the realised values of F. Let g be the implicit constraint equation defining the relationship between F and C, that is $g(C, F) = 0$. Then, n is the maximum cardinality of the set of real roots of g for each value of F:

$$n = \max_F |\{C \in \Re : g(C, F) = 0\}|. \qquad (B1-5)$$

There are several ways to calculate n. If $g(C, F) = 0$ is a polynomial in C, then an exact value for n can easily be obtained using Sturm's theorem for counting distinct real roots (Dorrie & Antin 1965). In the more general case, if $g(C, F) = 0$ is transcendental in C, then g can be approximated to arbitrary accuracy by Chebyshev polynomials, with real roots counted using companion matrix eigenvalue methods (Boyd 2013).

It is also possible to obtain an upper bound on the maximum state number. As we prove in Data S1, if F and C are both periodic in time, then the maximum state number is always finite, with

$$n \leq kb, \qquad (B1-6)$$

where k is the number of times F folds over itself in one period in the F-C plane, and b is the relative periodicity of F relative to C. The analytical bound essentially reflects how synchronised the observed and inferred climates are. Thus, even with *stable* dynamics characterised by periodic orbits, predictability of the community can vary strongly. Moreover, a simple functional form for $F(t)$ does not imply simple predictability in the inferred climate. The state number, which can be understood as a metric that characterises the *complexity* of those dynamics, is a valuable measure for predictability and the diversity of community responses possible.

the ball-in-cup landscapes used in ecosystem resilience/regime shift/alternate stable states research (e.g. Beisner *et al.* (2003); Scheffer & Carpenter (2003)) that also combine a parameter with a state variable. However, this diagram differs in that it shows the actual trajectory of the state variable over time, rather than the cost of taking different trajectories at a single point in time. That is, a community response diagram integrates the trajectories on a continually deforming ball-in-cup landscape and does not directly describe the stability or temporal dynamics of the community at any time point. As such, it is useful for addressing different questions than these other graphs, in particular questions of unstable or disequilibrium community responses to changing climate.
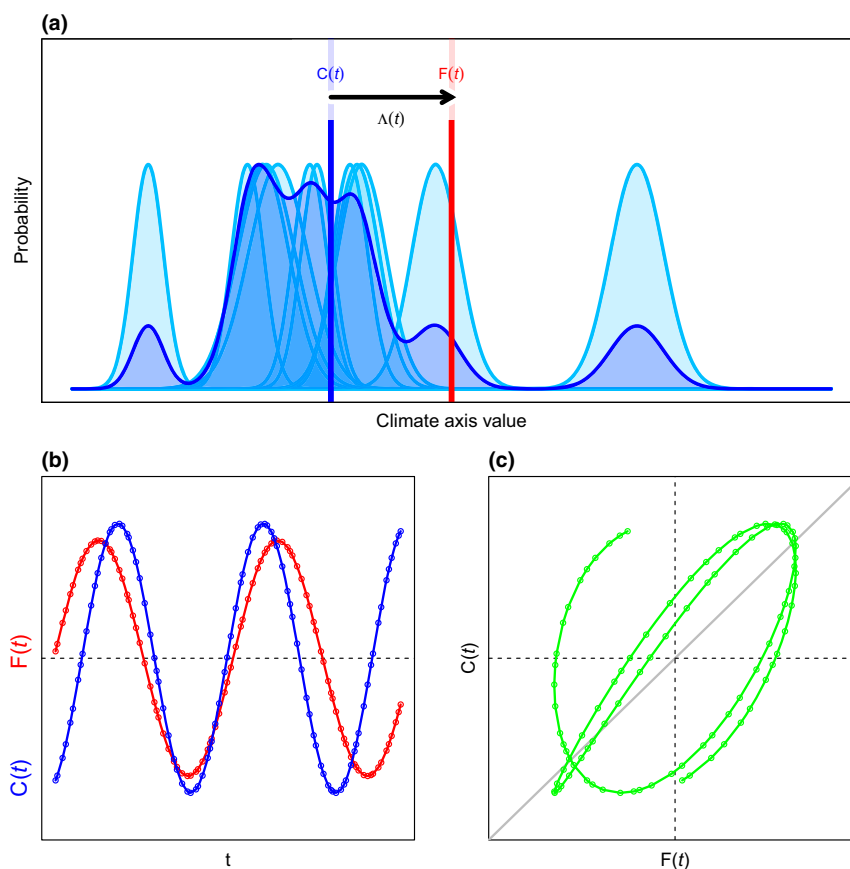
Figure 1 (a) Definition of community and climate terms. A community contains a set of resident species, each described by a different realised climate niche (cyan distributions) at time t. By overlapping these niches, a climate most consistent with the occurrence of these species (blue distribution) can be inferred and summarised by its expected value, defined as the community climate, $C(t)$ (vertical blue line). The community climate may differ from the observed climate at the location of the community, $F(t)$ (vertical red line). The difference between the community climate and the observed climate is defined as the community climate lag, $\Lambda(t)$. If the community is in equilibrium with climate and there are no lags, $\Lambda(t) = 0$, or $|\Lambda(t)| > 0$ otherwise. (b) An example of time series for $C(t)$ and $F(t)$. Values of zero are shown as a dashed horizontal line. (c) A community response diagram is a parametric plot of time series of $F(t)$ and $C(t)$. Data are replotted here from panel (b). Values of zero are shown as a dashed horizontal and vertical lines. The 1:1 no-lag expectation of $C(t) = F(t)$ is shown as a grey line.

By plotting the community's response as a function of the climate forcing, the continuum of lag hypotheses can be described and distinguished with two novel statistics. The first statistic is the mean absolute deviation, $\overline{|\Lambda|}$, which describes the average absolute difference between $C(t)$ and $F(t)$ over time (Fig. 2a). A value statistically indistinguishable from zero indicates no lag, and larger values indicate a lag (positive or negative). The second statistic is the maximum state number, n, which counts the maximum number of values of $C(t)$ that correspond to a single value of $F(t)$ (Fig. 2b). Considering the community response diagram as a curve in the F-C plane, n is the maximum number of intersection points of any vertical line. If there is only one value of $C(t)$ corresponding to each value of $F(t)$, then $n = 1$, and the community has dynamics that can always be predicted from knowledge of the current value of $F(t)$. If n becomes larger, then the community can have possible multiple states for a single observed climate. In these cases, it becomes increasingly less possible to predict the community's state with knowledge of only the observed climate. Thus, the maximum state number provides a simple

way to assess the limits to predictability for community dynamics.

A CONTINUUM OF LAG SCENARIOS ON A COMMUNITY RESPONSE DIAGRAM

There are several general scenarios for the coupling between climate change and community response that yield different $C(t)$ vs. $F(t)$ trajectories on a community response diagram (Fig. 3). Each of these scenarios also yields a different combination of values for the $\overline{|\Lambda|}$ and n statistics. Therefore, values of these statistics can be used to delineate hypotheses along the lag continuum.

The first scenario corresponds exactly to the no-lag hypothesis: in this case $\Lambda(t) = 0$, so $C(t) = F(t)$. This is equivalent to a straight-line segment with slope of 1 and intercept of 0 on the community response diagram for any possible observed climate $F(t)$ (Fig. 3a). In this scenario, $n = 1$ and $\overline{|\Lambda|} = 0$. Here, equality is statistically defined relative to natural variation, for example $\sigma(t)$.
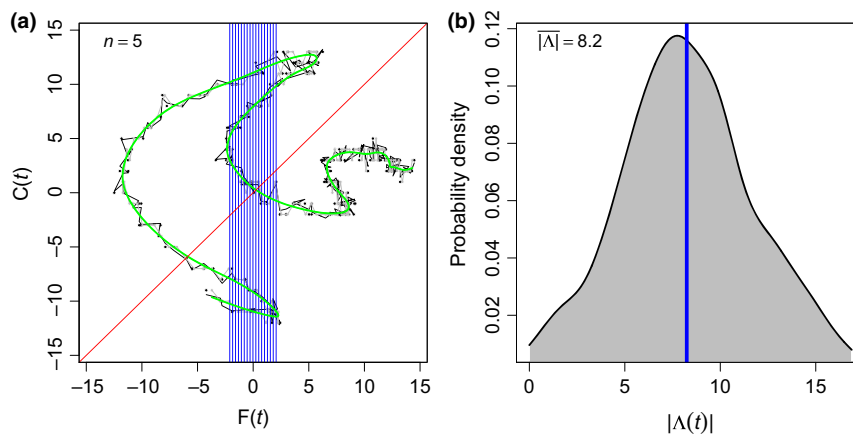
Figure 2 Definition of community response diagram statistics using an example data set. (a) A community's trajectory of observed climate $F(t)$ and the community response $C(t)$ is shown for original data (black curve), coarsened data (grey curve) and coarsened and smoothed data (blue curve). The 1 : 1 (no lag) expectation is shown as a diagonal red line. The maximum state number, n, indicates the largest number of unique values of $C(t)$ that correspond to any coarsened value of $F(t)$. It is calculated by intersecting a vertical line with the community's trajectory at all values of $F(t)$ (vertical blue lines). (b) The mean absolute deviation, $\overline{|\Lambda|}$, indicates the average difference between $C(t)$ and $F(t)$ across all times, with larger values indicating greater lags. The distribution of lags is shown as a grey envelope, and the statistic's value is shown as a vertical blue line.

The second scenario corresponds to a constant-relationship lag hypothesis. In this case, $|\Lambda(t)| \geq 0$ and $C(t) = f(F(t))$. Because f is a function, then there is always a single unique value of $C(t)$ corresponding to a unique value of $F(t)$. However, the opposite is not true: there may be multiple values of $F(t)$ that all correspond to the same value of $C(t)$. That is, the community's inferred climate is uniquely determined by the observed climate at any given time. This is equivalent to a single curve on the community response diagram that never crosses itself for any observed climate, so $n = 1$ and $\overline{|\Lambda|} > 0$ (Fig. 3b).

The third scenario corresponds to the constant-lag hypothesis. In this case, $|\Lambda(t)| = 0$ and $C(t) = \alpha F(t - \varphi)$ for some magnitude α and phase delay φ. If $F(t)$ is a periodic function, then this corresponds to a fold on the community response diagram, that is a scenario where $F(t)$ crosses over itself (Fig. 3b). In the case of a sinusoidal $F(t)$, the shape will be a single loop, with the elongation of the loop being related to the amount of lag (Fig. 3c). Such a scenario always has a value of $n = 2$ and $\overline{|\Lambda|} > 0$. However, for a linear $F(t)$ function, the shape will be a straight line with slope not necessarily equal to 1 and intercept not necessarily equal to 0. That scenario reduces to the constant-relationship conceptualisation and has $n = 1$ and $\overline{|\Lambda|} > 0$. In general, the presence of a fold or loop in the community response diagram indicates memory effects (hysteresis), such that the future state of the system depends on its past history (Katok & Hasselblatt 1997) (Fig. 3d). Systems with memory effects have path dependence. That is, the future dynamics of the community cannot be predicted only by knowing the current community state, but rather by also using the past state of the community. Larger values of $\overline{|\Lambda|}$ correspond to more memory effects.

The fourth scenario, alternate unstable states, is a generalised version of the third scenario, describing a community response diagram that contains multiple folds (Fig. 3e). At any given value of $F(t)$, the future state of the community

depends on its past state. If $F(t)$ is periodic, then the community response diagram will contain multiple loops corresponding to stable orbits. At any of the intersections between loops, determining which path the community takes will depend on knowledge of its past state. Alternatively, if the system has a stable orbit but has not yet reached it because of transient effects, then there may be large lags between $C(t)$ and $F(t)$ while the system settles to a steady state (Fig. 3c). These scenarios are all reflected in a value of $2 \leq n < \infty$ and $\overline{|\Lambda|} > 0$. Critically, these alternate states are not necessarily equivalent to the alternate stable states that have been previously studied (Beisner *et al.* 2003). They may not persist in time, and the community state is not necessarily attracted to them, although both scenarios are admissible. The key point here is that single values of the observed climate can lead to multiple values of the community state.

The fifth scenario, unpredictable dynamics, corresponds to a scenario where there are no stable orbits and a very large number of possible relationships between the observed climate and the composition of the community. Predicting the future state of the community is very difficult, because arbitrarily large changes in the community's future state can occur regardless of changes in values of the past community state or observed climate. In this case $n \to \infty$ and $\overline{|\Lambda|} > 0$. These dynamics can occur via chaos (Lorenz 1995), when the future community state is deterministic but very sensitive to variation in the present and past community state, where any state of the system is eventually reached from any other past state of the system, and where dynamical orbits are dense (Fig. 3f). Unpredictable dynamics can also occur when the future state of the community is not a deterministic response to any variable, as in the previous five scenarios, but rather is a stochastic response. In this case, $C(t)$ and $F(t)$ can become partially or completely uncorrelated, and a range of points in the community response diagram can become filled in (Fig. 3g). For example, random immigration and emigration of species from a regional species pool can yield fluctuations in $C(t)$ (Holyoak

et al. 2005)), while the climate system drives fluctuations in $F(t)$. Alternatively, $C(t)$ may be determined primarily by internal processes (e.g. species interactions, anthropogenic factors) rather than external climate-mediated processes, leading to a complete decoupling of $C(t)$ and $F(t)$. For example, many North American and European forests are thought to have
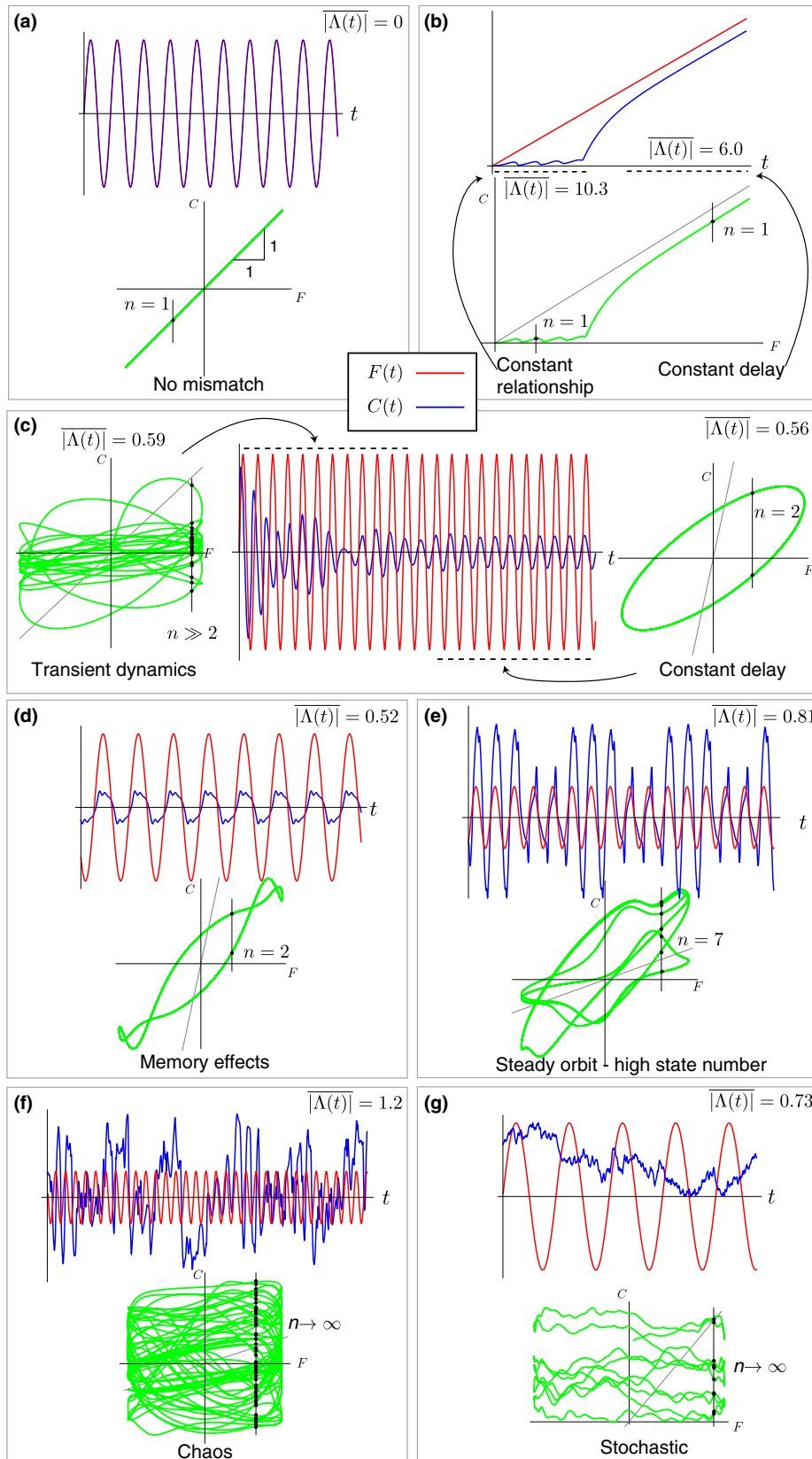
Figure 3 General classes of dynamics possible for a community's response to climate change. In each box, the time series shows an observed climate $F(t)$ (red lines) and a community response $C(t)$ (blue lines). Values of the state number n and the mean absolute deviation $\overline{|\Lambda|}$ are shown as insets for each example. (a) No-lag dynamics occur where the community climate closely matches the observed climate. This scenario can be detected when the community response diagram matches the 1 : 1 line. (b) Constant-relationship dynamics occur when the community response diagram is a function, that is has a unique value of $C(t)$ for every value of $F(t)$. (c) Constant delay dynamics occur when the community climate follows the observed climate with a fixed time delay. This scenario cannot be detected for a linear climate change but appears as a single loop for a sinusoidal climate change. Transient effects can also occur producing unpredictable dynamics with high n. (d) Memory effects occur when the community climate follows the observed climate with a variable delay and magnitude. This scenario can be detected via the presence of one or more crossing-back events that can also form loops when $F(t)$ is periodic. (e) Alternate unstable states occur when the community shows memory effects with multiple stacked loops, such that the state number is always greater than two. (f) Unpredictable dynamics can occur when n becomes infinite. Memory effects occur in this scenario as well. A scenario is shown here for chaos. (g) Unpredictable dynamics can also occur when the community response is uncorrelated with the observed climate, for example because of stochastic dynamics. All trajectories were generated from the model in Box 2 using parameter combinations described Table S1.

been managed for food production throughout the Holocene (Mason 2000; Abrams & Nowacki 2008), and many invasive species have colonised new regions due to enemy release (Keane & Crawley 2002), leading to geographic range shifts that are unrelated to climate change.

Each of these scenarios has different consequences for predictability in community ecology. The first two scenarios (no lag, constant-relationship lag) represent scenarios where prediction of future community states is readily possible. These scenarios have received the majority of study in community ecology, perhaps rightly. Nevertheless, the latter three scenarios are also conceptual possibilities. They challenge the assumptions of many research paradigms, because they imply there is no longer a simple or one-to-one relationship between climate conditions and community state. If constant-lag, alternative unstable states or unpredictable dynamics were to occur, then modelling a community's future state would be a challenge. With knowledge of only future observed climate, the task might be impossible; even with knowledge of the past observed climate and community state, the task might be very difficult. The consequence would be limited predictability in community ecology and shortened ecological forecast horizons (Petchey *et al.* 2015).

A SIMPLE ANALYTIC MODEL FOR LAGS IN COMMUNITY DYNAMICS

All of the scenarios along the lag continuum can arise within a simple differential equation model for community dynamics (Box 2). The model abstracts and summarises the community-scale effect of two species-scale processes: a tracking effect, in which communities try to restore themselves towards an optimal climate state, and a resistance effect, in which communities try to maintain their current (or past) composition. Temporal variation in the observed climate acts as a forcing for the model, while the interplay between the tracking and resistance processes determines the directionality and strength of the community's response.

The resistance and tracking effects are intended as proxies for a range of real ecological processes occurring for individual species. By abstracting these lower scale processes, we hope to gain general insights about possible dynamics at community scale. Both resistance and tracking must emerge from dispersal limitation, species interactions, environmental filtering or adaptation (Wisz *et al.* 2013; Singer *et al.* 2016; Zurell *et al.* 2016). For example, if some species in a regional pool

have a limited ability to disperse into a community, then tracking will be weaker. Similarly, resistance could be stronger if residents have advantages over invaders that must disperse in. If species interactions lead to established species persisting more easily than invader species can establish, then resistance will be stronger. Stronger environmental filtering could lead to stronger tracking by removing species with niches that yield low performance. Adaptation could yield both stronger tracking by shifting species' niches or stronger resistance by enabling species to maintain their niches. Our intent here is not to develop specific models that link these species-scale processes to community-scale effects, but rather to highlight how different types of such models would lead to different community outcomes. This exercise reveals several general principles of community dynamics. These conclusions are all mathematically true regardless of how the species-scale processes come together to yield a given set of restorative and tracking effects. First, linear climate change can only lead to no-lag, constant delay or constant-relationship scenarios regardless of all other model parameters, including the time delay Δt. However, under periodic climate change, all lag scenarios are possible. When the tracking and resistance effects are restorative, the system is characterised by transient dynamics towards a stable orbit. During the transient stage, n can become arbitrarily large, strongly limiting the ability to predict future states. The stronger the relative effect of the restorative effect, the longer the transient behaviour persists. However, $C(t)$ will eventually settle to a stable orbit with the same frequency as the observed climate $F(t)$. Depending on the exact form of the model, this steady orbit may constitute constant delay dynamics, but may also exhibit alternate unstable states with multiple loops. If the restorative effect pulls the system towards a fixed state with time delay (e.g. strong selection for a certain forest type regardless of climate), then along with the scenarios above, for different parameter regimes the system can also exhibit transient dynamics converging to alternate states with high n, as well as chaos.

The example community trajectories diagrammed in Figure 3 for each conceptual scenario correspond to dynamics predicted by this model for different parameter combinations. Specific parameter values are given in Table S1. However, the numeric values are less important than the general conclusion that the combination of restorative forces with time delay and climate tracking can lead to complex and widely varied dynamics, even in a simplified model. There are three conceptual implications arising from this modelling exercise that will

be relevant to all observational and theoretical studies of community dynamics.

First, the observation of a lag between the observed climate and the community response does not immediately indicate anything about the rules governing the system. Most scenarios show memory effects, so that knowledge of the past state of climate and community is needed to predict the system's future state. This result challenges the reliability of correlative methods for inferring the role of environmental drivers in community responses, because of the strong role of history on contemporary patterns (Dupouey *et al.* 2002; Willis *et al.* 2013). Nevertheless, for individual species, the success of species distribution models in predicting across space and time (Svenning & Sandel 2013) and the partial congruence of range limits and niche limits (Hargreaves *et al.* 2014; Lee-Yaw *et al.* 2016) suggest that simple no-lag approaches are viable. The major challenge will come in integrating species-scale predictions to community-scale responses where species-stacking approaches may fail because of interactions between species.

Second, small changes in a model's definitions can lead to qualitatively different types of ecological dynamics. This indicates that complex dynamics may become relevant in natural communities. Indeed, lagged dynamics with memory effects

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Box 2 A simple model of community dynamics

We propose an ordinary differential equation model for the dynamics of a community's state, $C(t)$. The model's formulation is general, but is operationalised here with linear functions to demonstrate the range of complex behaviour that can arise from simple model structure. The rate of change of $C(t)$ can be written as the sum of two processes as:

$$\frac{dC(t)}{dt} = -c_T T[\tau(t)] - c_R R[\rho(t)], \qquad (B2-1)$$

Here $T[\tau(t)]$ is a function describing how the community tracks a change $\tau(t)$ in its state relative to the observed climate at time t and $R[\rho(t)]$ is a function describing how the community resists a change $\rho(t)$ in its state at time t relative to a past observed climate. The coefficients $c_R \geq 0$ and $c_T \geq 0$ determine the relative importance of each effect. This model describes a forced delay differential equation, whose general properties and solutions have been explored in the mathematics and control theory literature (Sastry 2013).

The size of the tracking change, $\tau(t)$, can be defined as the linear difference between the observed climate and the community composition at time t:

$$\tau(t) = C(t) - F(t) = \Lambda(t). \qquad (B2-2)$$

We consider two possibilities for the resistance change, $\rho(t)$. One is to define resistance by the linear difference between the community composition at time t and the community composition based on a time delay, Δt:

$$\rho(t) = C(t) - C(t - \Delta t). \qquad (B2-3)$$

This models a scenario where the amount of restorative force is proportional to the difference between the community's past and present state, so that the system tends towards a past state (e.g. maintenance of an already-established forest type). Another is to use the difference between the community state at time $t - \Delta t$ and a optimal state C_0:

$$\rho(t) = C(t - \Delta t) - C_0, \qquad (B2-4)$$

which models a scenario where the system tends towards a fixed climate-independent optimum.

A simple proposal for the tracking function is a linear function:

$$T(\tau) = \tau. \qquad (B2-5)$$

where the response of a community to climate is directly proportional to the lag at that time.

Similarly, a simple resistance function can be proposed with a linear response, for example

$$R(\rho) = \rho. \qquad (B2-6)$$

or with a nonlinear response, as

$$R(\rho) = \rho^3. \qquad (B2-7)$$

Both resistance functions are odd and therefore yield responses that are restorative, in that they try to maintain the system in its current state. Another proposal is a nonlinear restorative function with multiple basins of attraction:

$$R(\rho) = \sin(\rho) \cdot \exp(-\rho^2). \qquad (B2-8)$$

This equation describes a situation where small to medium changes in system state lead to increasingly strong restorative responses, but where large changes lead to non-restorative responses.

The model also depends on the temporal trajectory of the observed climate $F(t)$. Here, we consider two simple example cases for climate change: a linearly increasing forcing with rate γ:

Box 2 (continued)

$$F_{\text{linear}}(t) = \gamma t \tag{B2 - 9}$$

and a periodic forcing with angular frequency ω:

$$F_{\text{periodic}}(t) = \sin(\omega t). \tag{B2 - 10}$$

First, consider the linear forcing. In the case that $c_R = 0$ (no resistance effects), Equation (B2-1) reduces to

$$\frac{dC(t)}{dt} = -c_T(C(t) - \gamma t) \tag{B2 - 11}$$

and has solution when $C(0) = 0$ of

$$C(t) = \gamma\left(t + \frac{e^{-c_T t} - 1}{c_T}\right) = F(t) - \frac{\gamma}{c_T}(1 - e^{-c_T t}). \tag{B2 - 12}$$

That is, the system is delayed by $\frac{\gamma}{c_T}(1 - e^{-c_T t})$. The second term rapidly decays over time, so the lag converges on a constant value as time increases. If $c_T = 0$, the delay become zero. Thus, only no-lag (Fig. 3a) or constant-lag (Fig. 3b) dynamics can occur.

If instead resistance does occur ($c_R > 0$), then Equation (B2-1) no longer has an exact solution. However, the system does respond with constant-relationship dynamics regardless of the choice of resistance function. Indeed, for any monotonic forcing function, this will be the case. For monotonic forcing, F and t are in a one-to-one relationship. Therefore, a given choice $F = F_0$ will correspond to a single time $t = t_0$. As $C(t)$ must be a function (emerging as the solution of a differential equation), fixing $t = t_0$ fixes $C = C_0 = C(t_0)$. This implies that even though $C(t)$ is not necessarily (in fact, usually not) a monotonic function, a given F corresponds to a single C, and thus, the community response diagram in the $F - C$ plane will be one-to-one, for which the state number is always $n = 1$ (Fig. 3b). The general implication is that only constant-lag, constant-relationship and no-lag dynamics are possible with linear climate change.

Next, consider the periodic forcing. The no-lag and constant delay hypotheses can both occur when there are no resistance effects ($c_R = 0$). The system reduces to

$$\frac{dC(t)}{dt} = -c_T(C(t) - \sin(\omega t)). \tag{B2 - 13}$$

In this case, the solution, assuming $C(0) = 0$, becomes

$$C(t) = \frac{c_T}{c_T^2 + \omega^2}\left[c_T \sin(\omega t) - \omega \sin\left(\frac{\pi}{2} - \omega t\right) + \omega e^{-c_T t}\right]. \tag{B2 - 14}$$

That is, the community response is proportional to the sum of the observed climate, a time delayed observed climate and a transient coefficient that decays rapidly over time (Fig. 3c). As the parameter c_T becomes large relative to ω, $C(t)$ converges exactly on $F(t)$ and the time lag disappears. That is, when $c_R = 0$, a small value of c_T corresponds to the constant-lag hypothesis, and a large value of c_T corresponds to the no-lag hypothesis.

Alternatively, when resistance effects also occur ($c_R > 0$), the type of dynamics depends on the size and form of the resistance (Table S1). For the simple lagged resistance (Equation (B2-3)), constant-lag and alternate state dynamics can occur, but are restricted to state number $n = 2$ (Fig. 3D). For the more complex restorative resistance change (Equation (B2-4)) and resistance functions (Equation (B2-7)), we find far more complex dynamics exhibited, including periodic states with state number $n \geq 2$ (Fig. 3E), as well as chaos in some parameter regimes (Fig. 3f).

We can also determine when (if ever) the system reaches a steady state, depending on the presence of resistance or tracking effects. As proved in Data S2, we can separate the community's dynamics into transient effects and steady states (except in the case of parameters leading to chaos). In the transient state, the system takes a trajectory that is highly influenced by initial conditions that can be difficult to predict. After the system settles to a steady state, $C(t)$ becomes a periodic function, and the community response diagram follows a fixed pattern that repeats over time. The duration of the transient increases with decreases in the ratio $\frac{c_T}{c_R}$, that is as resistance effects dominate tracking effects. The previous result holds except for where the tracking function restores towards a climate-independent state (Equation (B2-4)). In this case, are parameter regimes where the transient grows with time, even while the community is restoring towards a constant state. Thus, the system never obtains a fixed pattern that repeats over time and instead exhibits transient dynamics for all times that may have arbitrarily high n and $\overline{|\Lambda|}$.

and state numbers of $n \geq 2$ are easy to generate in the particular model we presented here.

Third, the frequency of climate change relative to the community response capacity is important in determining the type of dynamics that arise (Sastry 2013). When the observed

climate varies at much lower frequencies than the capacity of the community to respond, then the climate change is linear or effectively linear over the time period of interest, so no-lag or constant-relationship dynamics are likely to dominate (Williams *et al.* 2011). Similarly, when the observed climate varies

at much higher frequencies than the capacity of the tree community to respond, then unpredictable dynamics are likely to dominate unless the climate varies rapidly about a constant mean; then, the community may show limited response, as for example the case of *Populus tremuloides* – dominated tree communities that persist across glacial–interglacial transitions (Mitton & Grant 1996). Finally, when the observed climate varies at frequencies comparable to the community's response capability, then alternate unstable state or unpredictable dynamics may become important.

PRACTICAL CONCEPTUAL CONSIDERATIONS

A natural question arising from these conceptual and analytical arguments is as follows: which scenarios are likely to be found in the natural world? That is, is predictability achievable in practice, or not? The framework we have proposed could be applied to empirical data to answer this question. The scope of this article prevents presentation of such an analysis, so we instead focus on highlighting several issues that should be considered before implementing the framework.

Describing patterns of lags and testing lag hypotheses can be achieved by estimating $C(t)$ and $F(t)$ from data. If both time series are obtained from a finite number of empirical samples $\{C(t_i), F(t_i)\}$, then the easiest way to calculate both statistics is through approximation. Values of $\overline{|\Lambda|}$ can be obtained by averaging sampled values of $|C(t_i) - F(t_i)|$. Values of n can be calculated by linearly interpolation between successive values of $C(t_i)$ and $F(t_i)$ followed by application of line intersection methods. Measurement uncertainty or other noise arising in both $C(t)$ and $F(t)$ can be problematic when counting the maximum state number or determining whether a community response diagram contains loops. For example, suppose the community at two time points t_1 and t_2 has $F(t_1) = F(t_2)$, but $|C(t_1) - C(t_2)| < \max(\sigma(t_1), \sigma(t_2))$. In this case, the community takes two different states for a given observed climate, suggesting $n \geq 2$, but those states may not be sufficiently different to be confident that the difference is statistically significant. Regardless, the general qualitative implication is that estimation uncertainty and noise in time series can overestimate the maximum state number. Alternatively, low sampling resolution or a low number of points in a time series can lead to underestimation of n. Small loops or folds can be missed if they appear and disappear more rapidly than the sampling permits. We therefore recommend that community trajectories should be potentially rounded to the nearest multiple of $\sigma(t)$ and also smoothed before analysis (e.g. with cubic splines) (Fig. 2). We have implemented methods to calculate n and $\overline{|\Lambda|}$, taking into account statistical uncertainty in data, as R functions in Data S2.

Determining the underlying processes that have generated an empirical community response diagram is possible by fitting an analytic model to observed data for $C(t)$ and $F(t)$. There are several methods available to reconstruct a differential equation for $C(t)$ based on observations of $C(t)$ and $F(t)$ at different times. For example, generalised additive models with terms describing different effects can be fitted to numerical estimates of the first derivative of $C(t)$, providing a direct reconstruction of a differential equation (Ellner *et al.* 1997). It is also possible to estimate equation parameters from knowledge of the distribution of time intervals between extremes in a data set (Bezruchko *et al.* 2001). Alternatively, equation-free approaches for predicting the future state of a system based on its past state may be viable (Sugihara *et al.* 2012; Ye *et al.* 2015). However, all of these approaches tend to require data sampled at hundreds of different times, which may not be achievable for ecological data.

In either application, it may be difficult to distinguish among transient, chaotic and stochastic dynamics. For finite numbers of samples, these all lead to coarsely similar community response diagrams. When sampling of $C(t)$ and $F(t)$ is infrequent or includes measurement errors, it may be difficult to separate signal from noise in community response diagrams. Formal tests for distinguishing chaos from noise based on embedding of dynamical systems do exist (Gottwald & Melbourne 2004). Only with very long time series and precise measurements would it be possible to distinguish these scenarios in practice.

Lastly, it may be challenging to make unbiased measurements of $C(t)$. Because $C(t)$ depends on knowing the modal niche value for each species, any bias in species' estimating species niches may also propagate to community-scale statistics. Realised niche estimates based on contemporary geographic occurrences of species maybe particularly biased and themselves show lags (Jackson & Overpeck 2000; Soberón & Nakamura 2009), but provide the simplest method for calculating these statistics (Blonder *et al.* 2015).

PRACTICAL DATA CONSIDERATIONS

Finding data to infer community response diagrams remains a challenge. A representative sample of the community's composition is required to estimate $C(t)$. Time series of community dynamics are rare because of the long time scales and high efforts involved in this sampling. The best example is probably from the Park Grass Experiment in England, comprising dozens of censuses between 1856 and 2006 (Silvertown *et al.* 2006). On the other hand, the 50-ha forest dynamics plot at Barro Colorado Island (Panama) has been censused only seven times between 1980 and 2010 (Condit *et al.* 2012), and the macrophyte communities at Loch Leven (Scotland) have been censused only eight times between 1905 and 2008 (Dudley *et al.* 2012). Temporal extent is less important than number of time points: for example, it is possible to compare vegetation change on Chimborazo volcano (Ecuador) between 1802 and 2012, but measurements are only available at those two time points (Morueta-Holme *et al.* 2015). Other highly sampled time series, for example the Isle Royale (United States) wolf-moose data set (Vucetich & Peterson 2012), are oriented towards a single focal species rather than whole communities.

However, there are some systems where representative samples of communities at multiple time points are available. Microcosm studies of provide one possibility, for example protist communities (Petchey *et al.* 1999); similarly, metagenomics approaches are making community dynamics in microbial communities increasingly accessible (Faust *et al.*

2015). Alternatively, at longer time scales, paleoecological assemblage data sets may provide proxies for community dynamics. For example, fossil pollen assemblages for eastern North America are available for the last 21 Kyr at 500 year resolution, for example Maguire *et al.* (2016). Indeed, many of the studies that have calculated inferred climate time series using other approaches, for example Mosbrugger & Utescher (1997); Kühl *et al.* (2002); DeVictor *et al.* (2008); Bertrand *et al.* (2011), could be recast in terms of $C(t)$. In these cases, the challenge would be to make estimates of $F(t)$ for these communities that are independent of the community data. In the case of late Quaternary climate change, paleoclimate simulations based on general circulation models provide proxies for $F(t)$, but spatial and temporal resolution still remains coarse (Lorenz *et al.* 2016). For more contemporary time series, meteorological data may instead be available, for example Bertrand *et al.* (2011). Better understanding the limitations and potential of these various data sets, as well as actively collecting more time series community data, remains an ongoing but important monitoring challenge for ecology.

IMPLICATIONS AND SYNTHESIS

We showed that community response diagrams comprising plots of $F(t)$ and $C(t)$ provide methods to assess and understand how climate drives disequilibrium community states. By measuring lags with these community trajectories, and by calculating mean absolute deviations and maximum state numbers, we are able to provide approaches to assess the continuum of lag hypotheses, determine the limits to predictability and assess the importance of a community's past on its future.

The possibility of memory effects underscores the challenges present in predictive community ecology. Hysteresis is known to limit the ability of systems to return to an original stable state (Beisner *et al.* 2003; Folke *et al.* 2004), but our work now shows that unstable communities are also not guaranteed to return to the same state when the observed climate takes a previous value. This may provide a complementary explanation for why returning environments to historical conditions is unlikely to result in community shifts towards historical states: regime shifts can occur when the history of the community determines which future compositional state will be obtained (Scheffer & Carpenter 2003). Additionally, hysteresis suggests that commonly used space-for-time substitutions may not be appropriate, because the temporal dynamics of a system will depend on the past community state, while the spatial dynamics will not. Lastly, it also suggests that conservation efforts that take actions to reduce climate lags (such as assisted migration, rewilding or restoration of historical states) may potentially yield unexpected outcomes.

The niche axes comprising $C(t)$ and $F(t)$ can include any variables that mediate species response to environment. Here, we defined them in terms of climate variables. However, edaphic variables could also be important, given the close association between species occurrence and soil conditions. Soil legacies can persist for 10^3–10^4 years (Dupouey *et al.* 2002), and many species' distributions are very sensitive to soil conditions (Harrison 1999; Silvertown *et al.* 1999; Asner & Martin 2016), with community lags being driven by soil development (Kuneš *et al.* 2011).

The community response diagrams could also be recast in terms of functional traits, where $F(t)$ is an optimum trait value, and $C(t)$ is a community-weighted mean trait value (Garnier *et al.* 2004). Shifts and lags in trait–environment relationships (Kimberley *et al.* 2016; van der Sande *et al.* 2016) or skewness in trait distributions (Enquist *et al.* 2015) may be explainable using this approach. Remotely sensed community-weighted mean traits and remotely sensed climate data may be appropriate to explore this idea (e.g. Seddon *et al.* (2016)).

A next step towards more mechanistic understanding of community dynamics will be to couple the community-scale differential equation models to process-based models for individuals and populations of species in regional pools. By assessing the individual and combined effects of different types and strengths of dispersal limitation, species interactions and environmental filtering on community-scale patterns, it could become possible to identify the most likely drivers of each type of dynamics. Understanding the processes that lead to predictability and those that do not would help delineate when community ecology can hope to become more predictive (Fukami 2015) and when forecast horizons must remain small (Petchey *et al.* 2015).

CONCLUSION

We have explored some of the limits to predictability in community ecology using community response diagrams. The overall implication of this work is that predicting community response to past and near-future climate change will be difficult because of the diversity of possible dynamics. The no-lag hypothesis implicit in contemporary species distribution modelling represents a very narrow class of dynamics that may be successful at the scale of single species but not successful at the emergent community scale. The constant-lag and constant-relationship lag hypotheses of contemporary disequilibrium ecology and the extensions of species distribution modelling that incorporate dispersal limitation also represent a limited class of dynamics. There is evidence that some communities have sensitive responses to climate change (Ackerly 2003; Shuman *et al.* 2009; Nogué *et al.* 2013) and can exhibit regime shifts (Folke *et al.* 2004), whereas some others do not show evidence for this (Nowacki & Abrams 2015). Similarly, evidence for niche equilibrium at species scale is highly mixed (Veloz *et al.* 2012; Lee-Yaw *et al.* 2016). The possible existence of alternate unstable states and unpredictable dynamics should lead to careful consideration of whether extant approaches have oversimplified our perception of community dynamics.

Better delineating when and why responses to climate change will differ among communities should become a priority. Progress on predicting rather than explaining dynamics remains elusive and will require better understanding how processes such as species interactions and dispersal limitation determine the dynamical rules for community dynamics in response to climate change. We therefore suggest caution in our ability to make robust predictions about the future.

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AUTHOR CONTRIBUTIONS

BB conceived the project, carried out analyses and wrote the manuscript. BB and DM developed theory. All authors provided feedback and contributed to writing. The co-author list was ordered to reflect primary contributions from BB and DM and then ordered alphabetically by last name.

REFERENCES

Abrams, M.D. & Nowacki, G.J. (2008). Native Americans as active and passive promoters of mast and fruit trees in the eastern USA. *Holocene*, 18, 1123–1137.

Ackerly, D. (2003). Community assembly, niche conservatism, and adaptive evolution in changing environments. *Int. J. Plant Sci.*, 164, S165–S184.

Araújo, M.B. & Peterson, A.T. (2012). Uses and misuses of bioclimatic envelope modeling. *Ecology*, 93, 1527–1539.

Asner, G.P. & Martin, R.E. (2016). Convergent elevation trends in canopy chemical traits of tropical forests. *Glob. Change Biol.*, 22, 2216–2227.

Beisner, B.E., Haydon, D.T. & Cuddington, K. (2003). Alternative stable states in ecology. *Front. Ecol. Environ.*, 1, 376–382.

Bertrand, R., Lenoir, J., Piedallu, C., Riofrio-Dillon, G., de Ruffray, P., Vidal, C. *et al.* (2011). Changes in plant community composition lag behind climate warming in lowland forests. *Nature*, 479, 517–520.

Bezruchko, B.P., Karavaev, A.S., Ponomarenko, V.I. & Prokhorov, M.D. (2001). Reconstruction of time-delay systems from chaotic time series. *Phys. Rev. E*, 64, 056216.

Birks, H.H. & Ammann, B. (2000). Two terrestrial records of rapid climatic change during the glacial–Holocene transition (14,000–9,000 calendar years BP) from Europe. *Proc. Natl Acad. Sci.*, 97, 1390–1394.

Birks, H. & Birks, H.H. (2008). Biological responses to rapid climate change at the Younger Dryas—Holocene transition at Kråkenes, western Norway. *Holocene*, 18, 19–30.

Birks, H.J.B., Heiri, O., Seppä, H. & Bjune, A.E. (2010). Strengths and weaknesses of quantitative climate reconstructions based on late-quaternary biological proxies. *Open Ecol. J.*, 3, 68–110.

Blonder, B., Nogués-Bravo, D., Borregaard, M.K., Donoghue II, J.C., Jørgensen, P.M., Kraft, N.J.B. *et al.* (2015). Linking environmental filtering and disequilibrium to biogeography with a community climate framework. *Ecology*, 96, 972–985.

Boyd, J. (2013). Finding the zeros of a univariate equation: proxy Rootfinders, Chebyshev interpolation, and the companion matrix. *SIAM Rev.*, 55, 375–396.

ter Braak, C.J.F. & Prentice, I.C. (1988). A theory of gradient analysis. *Adv. Ecol. Res.*, 18, 271–317.

Brooks, S.J. & Birks, H. (2000). Chironomid-inferred late-glacial and early-Holocene mean July air temperatures for Kråkenes Lake, western Norway. *J. Paleolimnol.*, 23, 77–89.

Campbell, I.D. & McAndrews, J.H. (1993). Forest disequilibrium caused by rapid little ice age cooling. *Nature*, 366, 336–338.

Condit, R., Lao, S., Pérez, R., Dolins, S.B., Foster, R. & Hubbell, S. (2012). Barro Colorado Forest Census Plot Data (Version 2012). *Center for Tropical Forest Science Databases*.

Davis, M.B. (1981). Quaternary history and the stability of forest communities. In *Forest Succession: Concepts and Application*. (eds West, D.C., Shugart, H.H., Botkin, D.B.). Springer New York New York, NY, pp. 132–153.

Davis, M.B. (1984). Climatic instability, time, lags, and community disequilibrium. In: *Community Ecology* (eds. Diamond, J. & Case, T.J.). Harper & Row, New York, pp. 269–284.

De Frenne, P., Rodríguez-Sánchez, F., Coomes, D.A., Baeten, L., Verstraeten, G., Vellend, M. *et al.* (2013). Microclimate moderates plant responses to macroclimate warming. *Proc. Natl Acad. Sci.*, 110, 18561–18565.

DeVictor, V., Julliard, R., Couvet, D. & Jiguet, F. (2008). Birds are tracking climate warming, but not fast enough. *Proc. Biol. Sci.*, 275, 2743–2748.

Dorrie, H. & Antin, D. (1965). Sturm's Problem of the Number of Roots. In: *100 Great Problems of Elementary Mathematics: Their history and solution*. Dover Publications, New York, pp. 112–116.

Dudley, B., Gunn, I.D.M., Carvalho, L., Proctor, I., O'Hare, M.T., Murphy, K.J. *et al.* (2012). Changes in aquatic macrophyte communities in Loch Leven: evidence of recovery from eutrophication? *Hydrobiologia*, 681, 49–57.

Dullinger, S., Gattringer, A., Thuiller, W., Moser, D., Zimmermann, N.E., Guisan, A. *et al.* (2012). Extinction debt of high-mountain plants under twenty-first-century climate change. *Nat. Clim. Chang.*, 2, 619–622.

Dupouey, J.L., Dambrine, E., Laffite, J.D. & Moares, C. (2002). Irreversible impact of past land use on forest soils and biodiversity. *Ecology*, 83, 2978–2984.

Ellner, S.P., Kendall, B.E., Wood, S.N., McCauley, E. & Briggs, C.J. (1997). Inferring mechanism from time-series data: delay-differential equations. *Physica D*, 110, 182–194.

Enquist, B.J., Norberg, J., Bonser, S.P., Violle, C., Webb, C.T., Henderson, A. *et al.* (2015). Scaling from traits to ecosystems: developing a general trait driver theory via integrating trait-based and metabolic scaling theories. *Adv. Ecol. Res.*, 52, 249–318.

Eriksson, O. (1996). Regional dynamics of plants: a review of evidence for remnant, source-sink and metapopulations. *Oikos*, 77, 248–258.

Faust, K., Lahti, L., Gonze, D., de Vos, W.M. & Raes, J. (2015). Metagenomics meets time series analysis: unraveling microbial community dynamics. *Curr. Opin. Microbiol.*, 25, 56–66.

Folke, C., Carpenter, S., Walker, B., Scheffer, M., Elmqvist, T., Gunderson, L. *et al.* (2004). Regime shifts, resilience, and biodiversity in ecosystem management. *Annu. Rev. Ecol. Evol. Syst.*, 35, 557–581.

Fukami, T. (2015). Historical contingency in community assembly: integrating niches, species pools, and priority effects. *Annu. Rev. Ecol. Evol. Syst.*, 46, 1–23.

Fukami, T., Martijn Bezemer, T., Mortimer, S.R. & van der Putten, W.H. (2005). Species divergence and trait convergence in experimental plant community assembly. *Ecol. Lett.*, 8, 1283–1290.

Fukami, T., Dickie, I.A., Paula Wilkie, J., Paulus, B.C., Park, D., Roberts, A. *et al.* (2010). Assembly history dictates ecosystem functioning: evidence from wood decomposer communities. *Ecol. Lett.*, 13, 675–684.

Garnier, E., Cortez, J., Billès, G., Navas, M.-L., Roumet, C., Debussche, M. *et al.* (2004). Plant functional markers capture ecosystem properties during secondary succession. *Ecology*, 85, 2630–2637.

Gasse, F., Juggins, S. & Khelifa, L.B. (1995). Diatom-based transfer functions for inferring past hydrochemical characteristics of African lakes. *Palaeogeogr. Palaeoclimatol. Palaeoecol.*, 117, 31–54.

Gottwald, G.A. & Melbourne, I. (2004). A new test for chaos in deterministic systems. *Proc. R. Soc. Lon. A*, 460, 603–611.

Guiot, J., Pons, A., de Beaulieu, J.L. & Reille, M. (1989). A 140,000-year continental climate reconstruction from two European pollen records. *Nature*, 338, 309–313.

Guisan, A. & Thuiller, W. (2005). Predicting species distribution: offering more than simple habitat models. *Ecol. Lett.*, 8, 993–1009.

Hampe, A. & Petit, R.J. (2005). Conserving biodiversity under climate change: the rear edge matters. *Ecol. Lett.*, 8, 461–467.

Harbert, R.S. & Nixon, K.C. (2015). Climate reconstruction analysis using coexistence likelihood estimation (CRACLE): a method for the estimation of climate using vegetation. *Am. J. Bot.*, 102, 1277–1289.

Hargreaves, A.L., Samis, K.E. & Eckert, C.G. (2014). Are species range limits simply niche limits writ large? A review of transplant experiments beyond the range. *Am. Nat.*, 183, 157–173.

Harrison, S. (1999). Local and regional diversity in a patchy landscape: native, alien, and endemic herbs on serpentine. *Ecology*, 80, 70–80.

Hoffmann, A.A. & Sgro, C.M. (2011). Climate change and evolutionary adaptation. *Nature*, 470, 479–485.

Holt, R.D. (2009). Bringing the Hutchinsonian niche into the 21st century: ecological and evolutionary perspectives. *Proc. Natl Acad. Sci.*, 106, 19659–19665.

Holyoak, M., Leibold, M.A. & Holt, R.D. (2005). *Metacommunities: Spatial Dynamics and Ecological Communities*. University of Chicago Press, Chicago, IL.

von Humboldt, A. & Bonpland, A. (1807. *Essay on the Geography of Plants*. (ed. Jackson, ST). University of Chicago Press, Chicago, IL.

Jackson, S.T. & Overpeck, J.T. (2000). Responses of plant populations and communities to environmental changes of the late Quaternary. *Paleobiology*, 26, 194–220.

Jackson, S.T. & Sax, D.F. (2010). Balancing biodiversity in a changing environment: extinction debt, immigration credit and species turnover. *Trends Ecol. Evol.*, 25, 153–160.

Jordan, G.J. (2011). A critical framework for the assessment of biological palaeoproxies: predicting past climate and levels of atmospheric CO_2 from fossil leaves. *New Phytol.*, 192, 29–44.

Katok, A. & Hasselblatt, B. (1997). *Introduction to the Modern Theory of Dynamical Systems*. Cambridge University Press, Cambridge.

Keane, R.M. & Crawley, M.J. (2002). Exotic plant invasions and the enemy release hypothesis. *Trends Ecol. Evol.*, 17, 164–170.

Kimberley, A., Blackburn, G.A., Whyatt, J.D. & Smart, S.M. (2016). How well is current plant trait composition predicted by modern and historical forest spatial configuration? *Ecography*, 39, 67–76.

Kühl, N., Gebhardt, C., Litt, T. & Hense, A. (2002). Probability density functions as botanical-climatological transfer functions for climate reconstruction. *Quatern. Res.*, 58, 381–392.

Kuneš, P., Odgaard, B.V. & Gaillard, M.-J. (2011). Soil phosphorus as a control of productivity and openness in temperate interglacial forest ecosystems. *J. Biogeogr.*, 38, 2150–2164.

La Sorte, F.A. & Jetz, W. (2012). Tracking of climatic niche boundaries under recent climate change. *J. Anim. Ecol.*, 81, 914–925.

Lee-Yaw, J.A., Kharouba, H.M., Bontrager, M., Mahony, C., Csergő, A.M., Noreen, A.M.E. *et al.* (2016). A synthesis of transplant experiments and ecological niche models suggests that range limits are often niche limits. *Ecol. Lett.*, 19, 710–722.

Lenoir, J., Graae, B.J., Aarrestad, P.A., Alsos, I.G., Armbruster, W.S., Austrheim, G. *et al.* (2013). Local temperatures inferred from plant communities suggest strong spatial buffering of climate warming across Northern Europe. *Glob. Change Biol.*, 19, 1470–1481.

Lorenz, E.N. (1995). *The Essence of Chaos*. University of Washington Press, Seattle.

Lorenz, D.J., Nieto-Lugilde, D., Blois, J.L., Fitzpatrick, M.C. & Williams, J.W. (2016). Downscaled and debiased climate simulations for North America from 21,000 years ago to 2100AD. *Sci. Data*, 3, 160048.

Maguire, K.C., Nieto-Lugilde, D., Blois, J.L., Fitzpatrick, M.C., Williams, J.W., Ferrier, S. *et al.* (2016). Controlled comparison of species- and community-level models across novel climates and communities. *Proc. R. Soc. Lon. B.*, 283, 20152817.

Maiorano, L., Cheddadi, R., Zimmermann, N.E., Pellissier, L., Petitpierre, B., Pottier, J. *et al.* (2013). Building the niche through time: using 13,000 years of data to predict the effects of climate change on three tree species in Europe. *Glob. Ecol. Biogeogr.*, 22, 302–317.

Mason, S. (2000). Fire and Mesolithic subsistence—managing oaks for acorns in northwest Europe? *Palaeogeogr. Palaeoclimatol. Palaeoecol.*, 164, 139–150.

Mitton, J.B. & Grant, M.C. (1996). Genetic variation and the natural history of quaking aspen. *Bioscience*, 46, 25–31.

Morueta-Holme, N., Engemann, K., Sandoval-Acuña, P., Jonas, J.D., Segnitz, R.M. & Svenning, J.-C. (2015). Strong upslope shifts in Chimborazo's vegetation over two centuries since Humboldt. *Proc. Natl Acad. Sci.*, 112, 12741–12745.

Mosbrugger, V. & Utescher, T. (1997). The coexistence approach — a method for quantitative reconstructions of Tertiary terrestrial palaeoclimate data using plant fossils. *Palaeogeogr. Palaeoclimatol. Palaeoecol.*, 134, 61–86.

Nicotra, A.B., Atkin, O.K., Bonser, S.P., Davidson, A.M., Finnegan, E.J., Mathesius, U. *et al.* (2010). Plant phenotypic plasticity in a changing climate. *Trends Plant Sci.*, 15, 684–692.

Nogué, S., de Nascimento, L., Fernández-Palacios, J.M., Whittaker, R.J. & Willis, K.J. (2013). The ancient forests of La Gomera, Canary Islands, and their sensitivity to environmental change. *J. Ecol.*, 101, 368–377.

Normand, S., Ricklefs, R.E., Skov, F., Bladt, J., Tackenberg, O. & Svenning, J.-C. (2011). Postglacial migration supplements climate in determining plant species ranges in Europe. *Proc. R. Soc. B.*, 278, 3644–3653.

Nowacki, G.J. & Abrams, M.D. (2015). Is climate an important driver of post-European vegetation change in the Eastern United States? *Glob. Change Biol.*, 21, 314–334.

Petchey, O.L., McPhearson, P.T., Casey, T.M. & Morin, P.J. (1999). Environmental warming alters food-web structure and ecosystem function. *Nature*, 402, 69–72.

Petchey, O.L., Pontarp, M., Massie, T.M., Kéfi, S., Ozgul, A., Weilenmann, M. *et al.* (2015). The ecological forecast horizon, and examples of its uses and determinants. *Ecol. Lett.*, 18, 597–611.

Peterson, A.T. (2011). *Ecological Niches and Geographic Distributions (MPB-49)*. Princeton University Press, Princeton.

Pross, J., Klotz, S. & Mosbrugger, V. (2000). Reconstructing palaeotemperatures for the Early and Middle Pleistocene using the mutual climatic range method based on plant fossils. *Quatern. Sci. Rev.*, 19, 1785–1799.

van der Sande, M.T., Arets, E.J.M.M., Peña-Claros, M., de Avila, A.L., Roopsind, A., Mazzei, L. *et al.* (2016). Old-growth neotropical forests are shifting in species and trait composition. *Ecol. Monogr.*, 86, 228–243.

Sastry, S.S. (2013). *Nonlinear Systems: Analysis, Stability, and Control*. Springer Science & Business Media, Springer-Verlag, New York, NY.

Scheffer, M. & Carpenter, S.R. (2003). Catastrophic regime shifts in ecosystems: linking theory to observation. *Trends Ecol. Evol.*, 18, 648–656.

Schöb, C., Butterfield, B.J. & Pugnaire, F.I. (2012). Foundation species influence trait-based community assembly. *New Phytol.*, 196, 824–834.

Seddon, A.W., Macias-Fauria, M. & Willis, K.J. (2015). Climate and abrupt vegetation change in Northern Europe since the last deglaciation. *Holocene*, 25, 25–36.

Seddon, A.W., Macias-Fauria, M., Long, P.R., Benz, D. & Willis, K.J. (2016). Sensitivity of global terrestrial ecosystems to climate variability. *Nature*, 531, 229–232.

Shuman, B.N., Newby, P. & Donnelly, J.P. (2009). Abrupt climate change as an important agent of ecological change in the Northeast U.S. throughout the past 15,000 years. *Quatern. Sci. Rev.*, 28, 1693–1709.

Silvertown, J., Dodd, M.E., Gowing, D.J.G. & Mountford, J.O. (1999). Hydrologically defined niches reveal a basis for species richness in plant communities. *Nature*, 400, 61–63.

Silvertown, J., Poulton, P., Johnston, E., Edwards, G., Heard, M. & Biss, P.M. (2006). The Park Grass Experiment 1856–2006: its contribution to ecology. *J. Ecol.*, 94, 801–814.

Singer, A., Johst, K., Banitz, T., Fowler, M.S., Groeneveld, J., Gutiérrez, A.G. *et al.* (2016). Community dynamics under environmental change:

how can next generation mechanistic models improve projections of species distributions? *Ecol. Model.*, 326, 63–74.

Soberón, J. & Nakamura, M. (2009). Niches and distributional areas: concepts, methods, and assumptions. *Proc. Natl Acad. Sci.*, 106, 19644–19650.

Sugihara, G., May, R., Ye, H., Hsieh, C.-H., Deyle, E., Fogarty, M. *et al.* (2012). Detecting causality in complex ecosystems. *Science*, 338, 496–500.

Svenning, J.-C. & Sandel, B. (2013). Disequilibrium vegetation dynamics under future climate change. *Am. J. Bot.*, 100, 1266–1286.

Svenning, J.-C. & Skov, F. (2007). Could the tree diversity pattern in Europe be generated by postglacial dispersal limitation? *Ecol. Lett.*, 10, 453–460.

Svenning, J.-C., Eiserhardt, W.L., Normand, S., Ordonez, A. & Sandel, B. (2015). The influence of paleoclimate on present-day patterns in biodiversity and ecosystems. *Annu. Rev. Ecol. Evol. Syst.*, 46, 551–572.

Veloz, S.D., Williams, J.W., Blois, J.L., He, F., Otto-Bliesner, B. & Liu, Z. (2012). No-analog climates and shifting realized niches during the late quaternary: implications for 21st-century predictions by species distribution models. *Glob. Change Biol.*, 18, 1698–1713.

Vucetich, J. & Peterson, R. (2012). The population biology of Isle Royale wolves and moose: an overview. Available at: http://www.isleroyale wolf.org/data/data/home.html. Last accessed 10 January 2017.

Webb, T. (1986). Is vegetation in equilibrium with climate? How to interpret late-Quaternary pollen data. *Vegetatio*, 67, 75–91.

Whittaker, R.H. (1967). Gradient analysis of vegetation. *Biol. Rev.*, 42, 207–264.

Williams, J.W., Blois, J.L. & Shuman, B.N. (2011). Extrinsic and intrinsic forcing of abrupt ecological change: case studies from the late Quaternary. *J. Ecol.*, 99, 664–677.

Willis, K.J., Bennett, K.D., Burrough, S.L., Macias-Fauria, M. & Tovar, C. (2013). Determining the response of African biota to climate change: using the past to model the future. *Philos. Trans. R Soc. Lond. B.*, 368, 20120491.

Wisz, M.S., Pottier, J., Kissling, W.D., Pellissier, L., Lenoir, J., Damgaard, C.F. *et al.* (2013). The role of biotic interactions in shaping distributions and realised assemblages of species: implications for species distribution modelling. *Biol. Rev.*, 88, 15–30.

Ye, H., Beamish, R.J., Glaser, S.M., Grant, S.C.H., Hsieh, C.-H., Richards, L.J. *et al.* (2015). Equation-free mechanistic ecosystem forecasting using empirical dynamic modeling. *Proc. Natl Acad. Sci.*, 112, E1569–E1576.

Zurell, D., Thuiller, W., Pagel, J., S Cabral, J., Münkemüller, T., Gravel, D. *et al.* (2016). Benchmarking novel approaches for modelling species range dynamics. *Glob. Change Biol.*, 22, 2651–2664.

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